

27TH INTERNATIONAL LASER RADAR CONFERENCE

CLOUD-AEROSOL INTERACTIONS: RETRIEVING AEROSOL ÅNGSTRÖM EXPONENTS FROM CALIPSO MEASUREMENTS OF OPAQUE WATER CLOUDS



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CALIPSO

27th International Laser Radar Conference (ILRC)

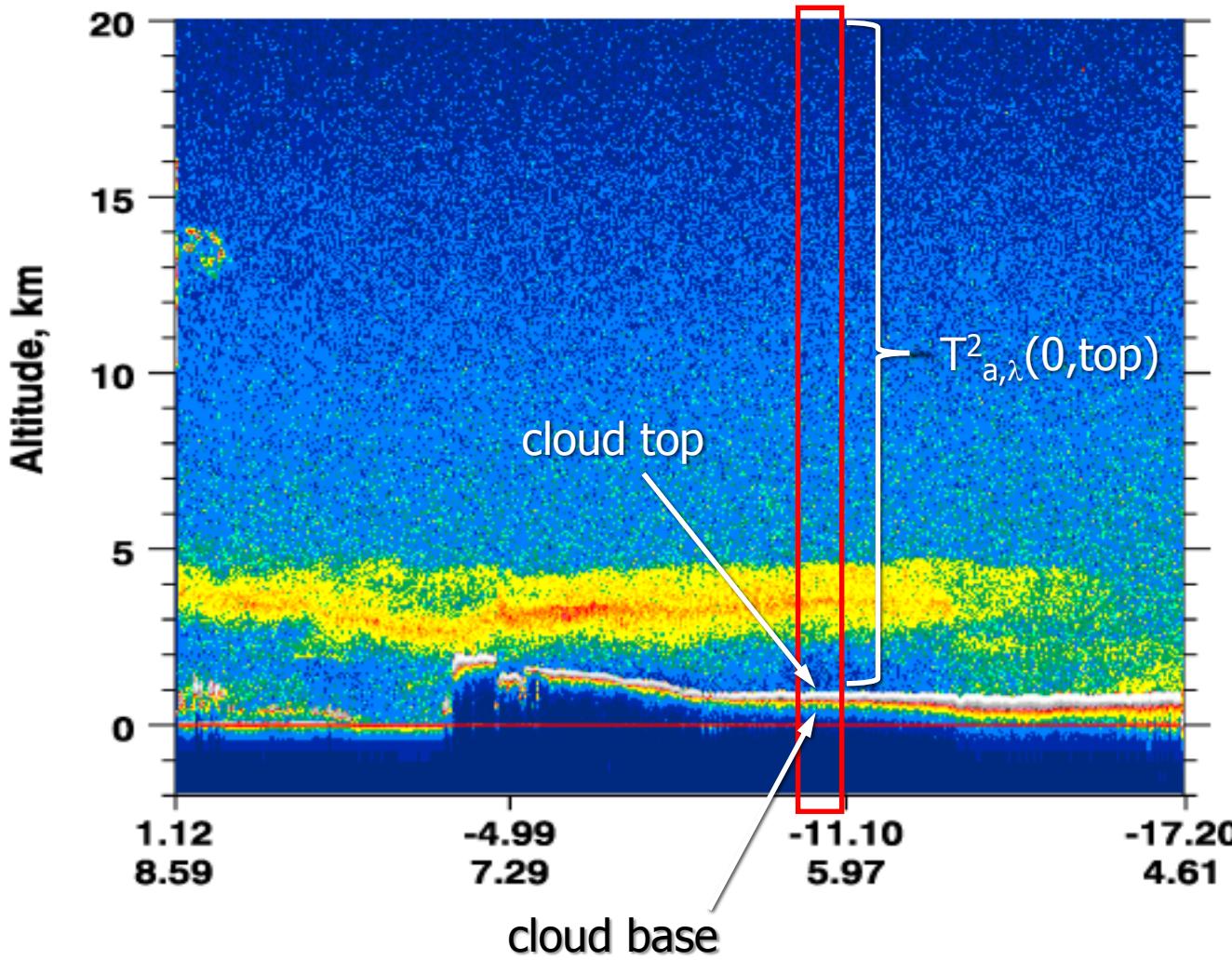
7 July 2015

Mark Vaughan (1)





HU ET AL., 2007: RETRIEVING OPTICAL DEPTHS AND LIDAR RATIOS FOR TRANSPARENT LAYERS ABOVE OPAQUE WATER CLOUDS



cloud integrated attenuated backscatter

$$\begin{aligned}\gamma'_{c,\lambda} &= T^2_{a,\lambda}(0, \text{CloudTop}) \int_{\text{CloudTop}}^{\text{CloudBase}} B_\lambda(z) dz \\ &= T^2_{a,\lambda}(0, \text{CloudTop}) \left(\frac{1}{2 \eta S_c} \right)\end{aligned}$$

Platt 1973

multiple scattering factor computed from measured in-cloud volume depolarization

$$\eta_{c,532} = \left(\frac{1 - \delta_{v,532}}{1 + \delta_{v,532}} \right)^2$$

above-cloud optical depth retrieved using a (known!) water cloud lidar ratio, S_c

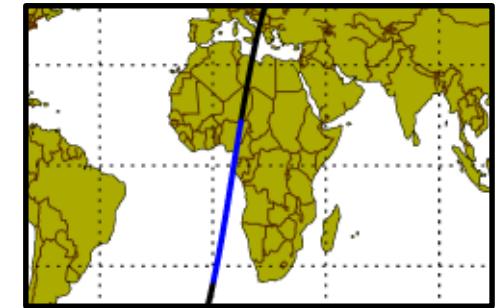
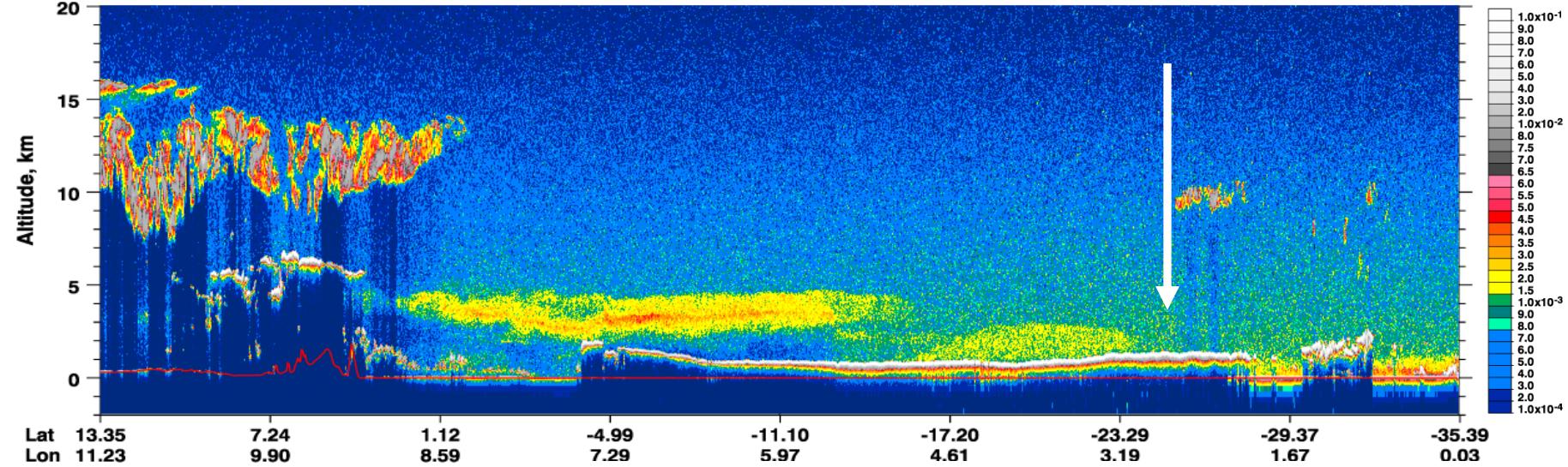
$$\begin{aligned}\tau_{a,532} &= -\frac{1}{2} \ln(2 S_c \gamma'_{c,532} \eta_{c,532}) \\ &= -\frac{1}{2} \ln \left(2 S_c \gamma'_{c,532} \left(\frac{1 - \delta_{v,532}}{1 + \delta_{v,532}} \right)^2 \right)\end{aligned}$$





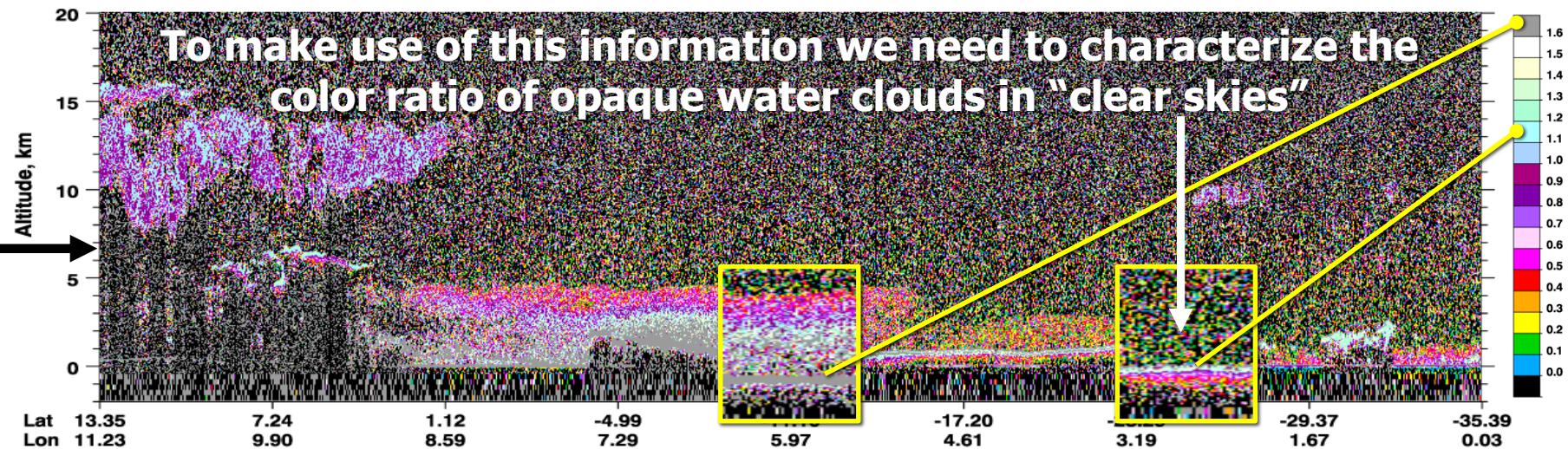
1064 nm DATA PROVIDES ADDITIONAL INFO

532 nm Total Attenuated Backscatter, $\text{km}^{-1} \text{sr}^{-1}$ UTC: 2008-08-04 01:08:15.3 to 2008-08-04 01:21:44.0 Version: 3.01 Nominal Nighttime



Attenuated Color Ratio, 1064nm/532nm UTC: 2008-08-04 01:08:15.3 to 2008-08-04 01:21:44.0 Version: 3.01 Nominal Nighttime

$$\chi'(z) = \frac{\beta'_{1064}(z)}{\beta'_{532}(z)}$$



To make use of this information we need to characterize the color ratio of opaque water clouds in "clear skies"

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LAYER-INTEGRATED ATTENUATED BACKSCATTER COLOR RATIO

$$\chi'_{\text{aerosol above}} = \frac{T_{a,1064}^2(0, \text{top}) \int_{\text{top}}^{\text{base}} B_{1064}(z) dz}{T_{a,532}^2(0, \text{top}) \int_{\text{top}}^{\text{base}} B_{532}(z) dz} = \frac{T_{a,1064}^2(0, \text{top})}{T_{a,532}^2(0, \text{top})} \chi'_{\text{clear}}$$

\$\chi'_{\text{aerosol above}} = \$T_{a,1064}^2(0, \text{top}) \int_{\text{top}}^{\text{base}} B_{1064}(z) dz\$ = \$T_{a,1064}^2(0, \text{top})\$ / \$T_{a,532}^2(0, \text{top})\$ \$\chi'_{\text{clear}}
↓ measured ↓ retrieved ↓ "characterized"

ONE EQUATION, ONE UNKNOWN

$$T_{a,1064}^2(0, \text{top}) = T_{a,532}^2(0, \text{top}) \frac{\chi'_{\text{aerosol above}}}{\chi'_{\text{clear}}}$$

$$\tau_{1064} = -0.5 \ln \left(T_{a,532}^2(0, \text{top}) \frac{\chi'_{\text{aerosol above}}}{\chi'_{\text{clear}}} \right)$$

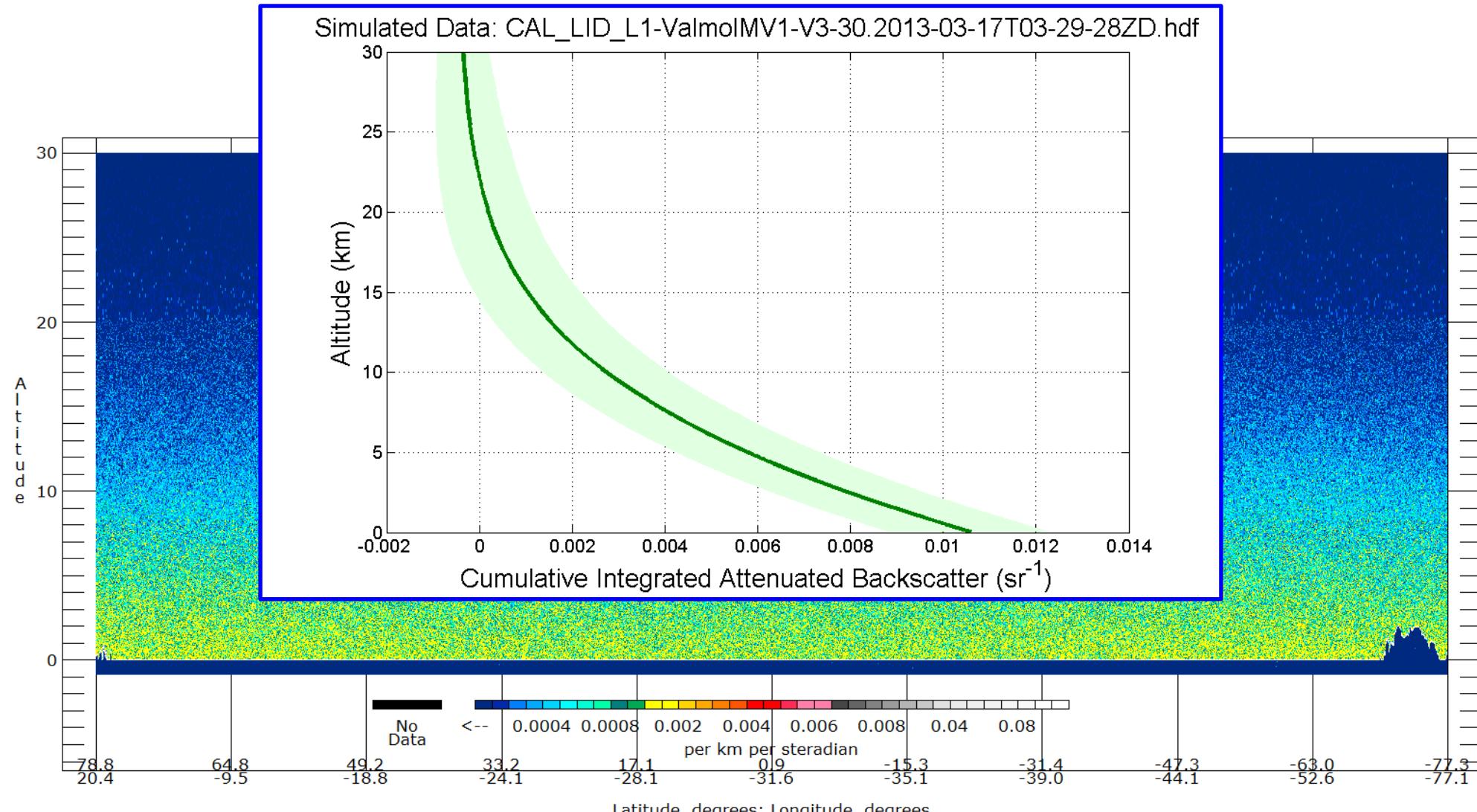
AEROSOL ÅNGSTRÖM EXPONENT

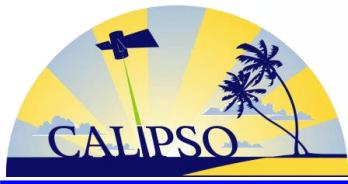
$$\alpha = - \frac{\log\left(\frac{\tau_{1064}}{\tau_{532}}\right)}{\log\left(\frac{1064}{532}\right)}$$



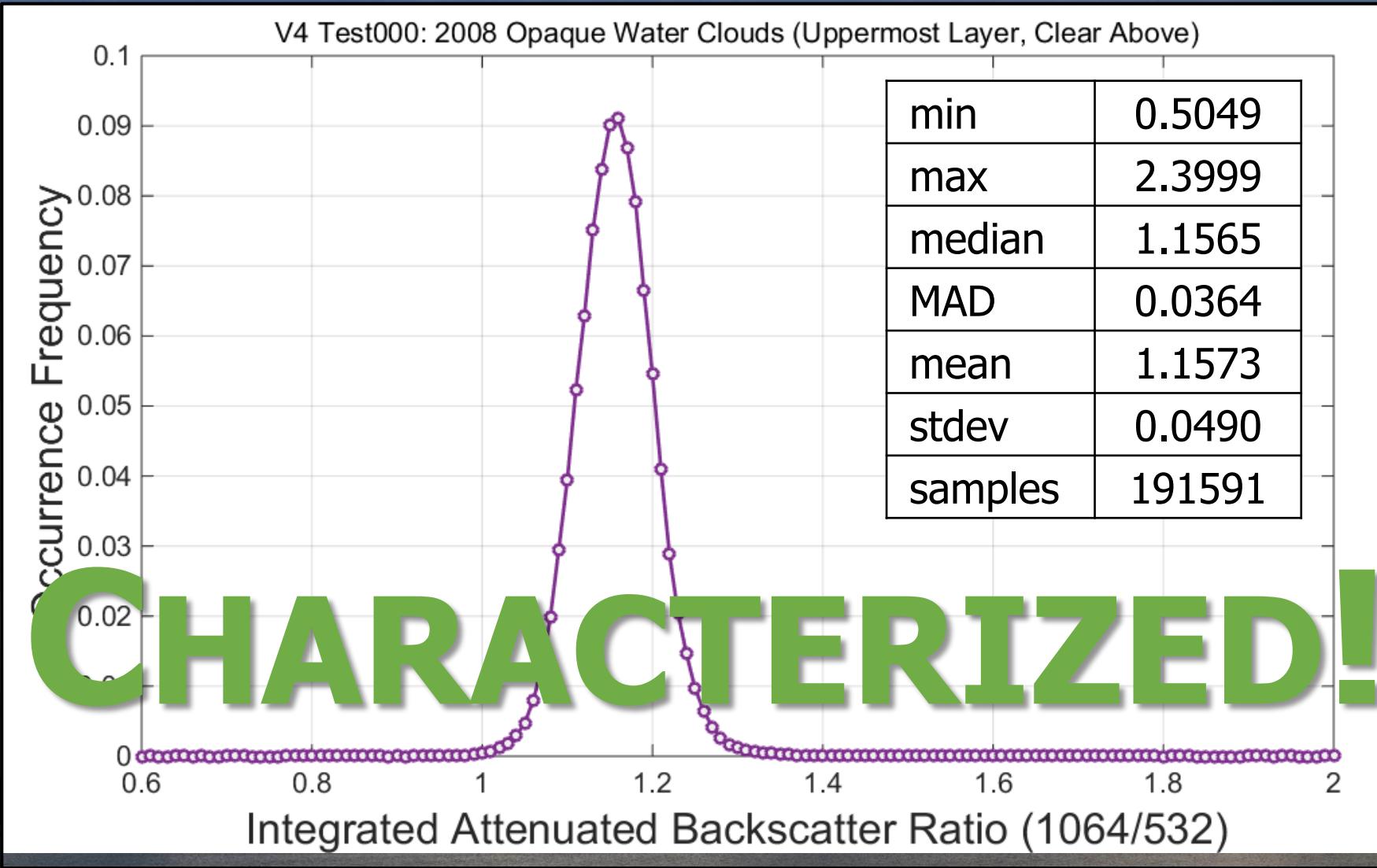


WHEN Do WE HAVE "CLEAR AIR" ABOVE?





WATER CLOUD COLOR RATIOS MEASURED BY CALIOP



LIU ET AL., 2015: EVALUATION OF CALIOP 532 nm AOD OVER OPAQUE WATER CLOUDS

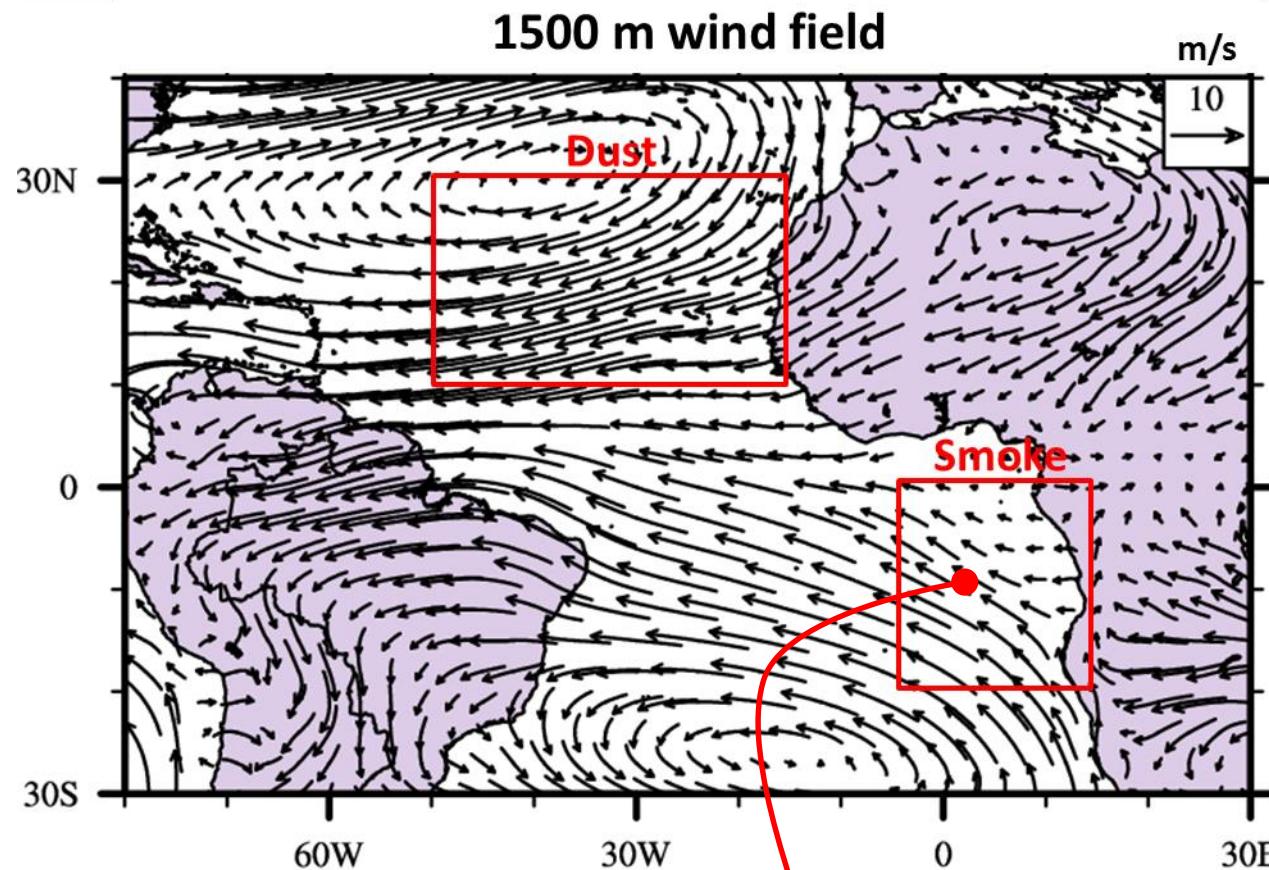


Figure 1: Spatial domains analyzed (red boxes) and wind fields (arrows) from ECMWF data for July and August from 2007 to 2012. The northern region (10°N - 30°N , 50°W - 15°W) is along the Saharan dust transport pathway over the tropical North Atlantic, while **the southern region (20°S - 0° , 5°W - 15°E) is along the smoke transport pathway over the tropical South Atlantic.**

Smoke Lidar Ratios: 2007-2012

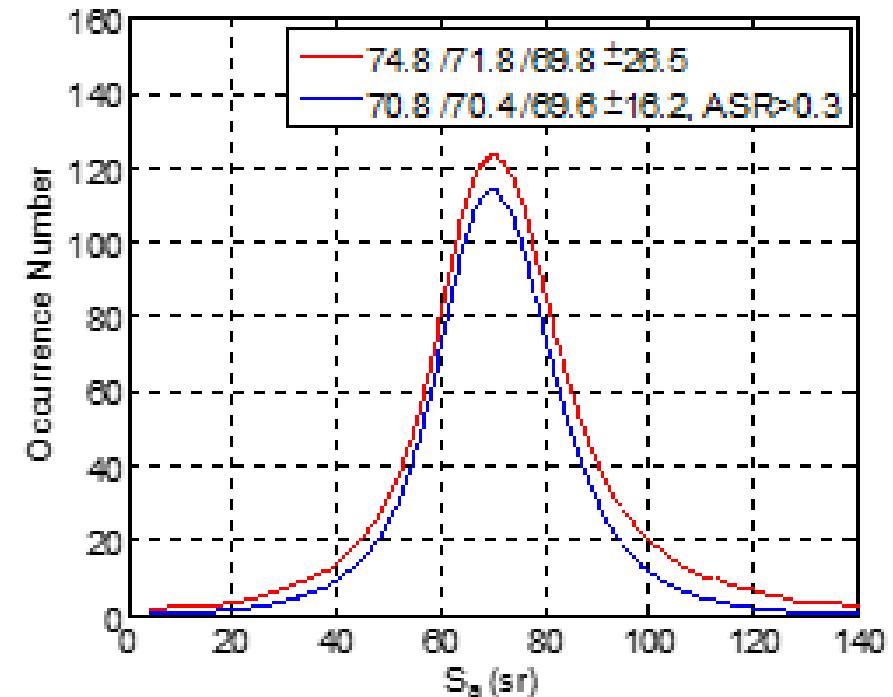
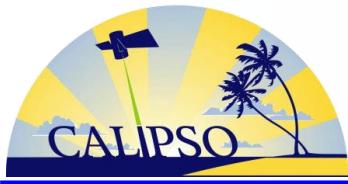
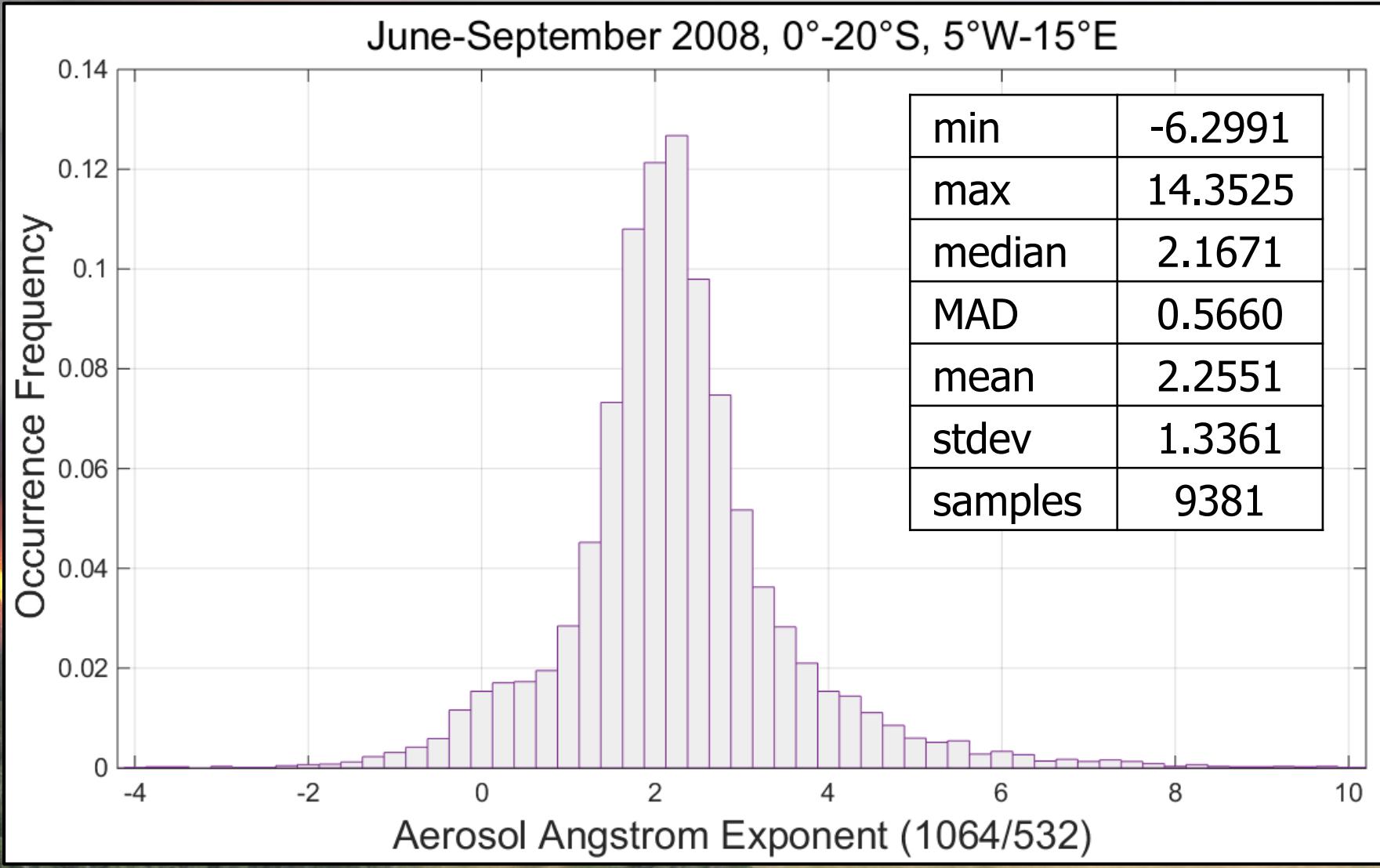


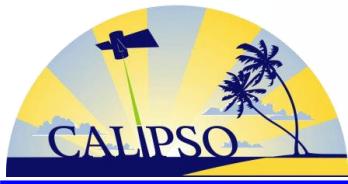
Figure 10(c): The red curve show result for all data, while the blue curve includes only data with mean scattering ratios above 0.2. The numbers in the legends of are mean/median/mode \pm standard deviation.





CALIOP ÅNGSTRÖM EXPONENTS FOR SMOKE





CALIOP LIDAR RATIOS FOR SMOKE

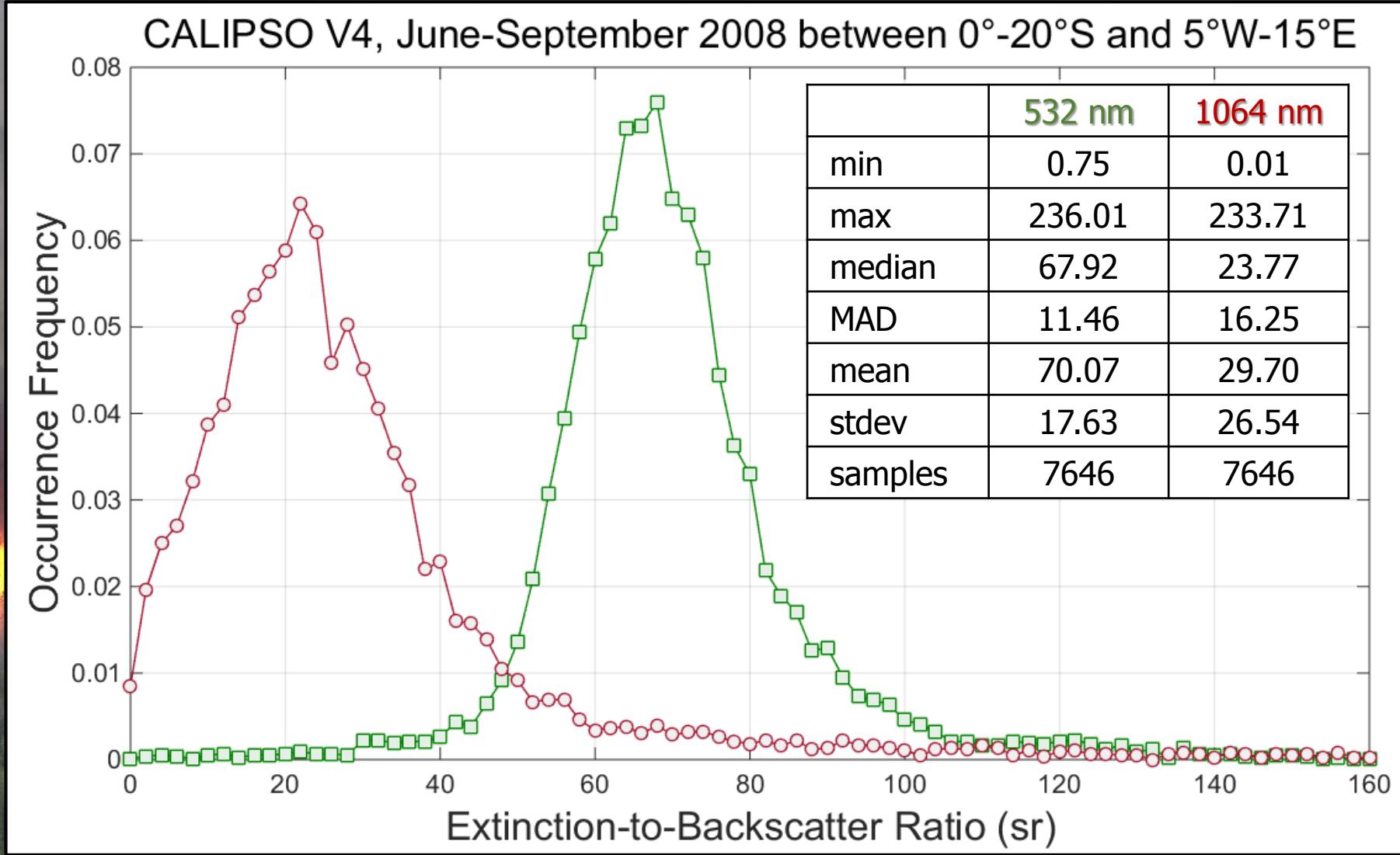


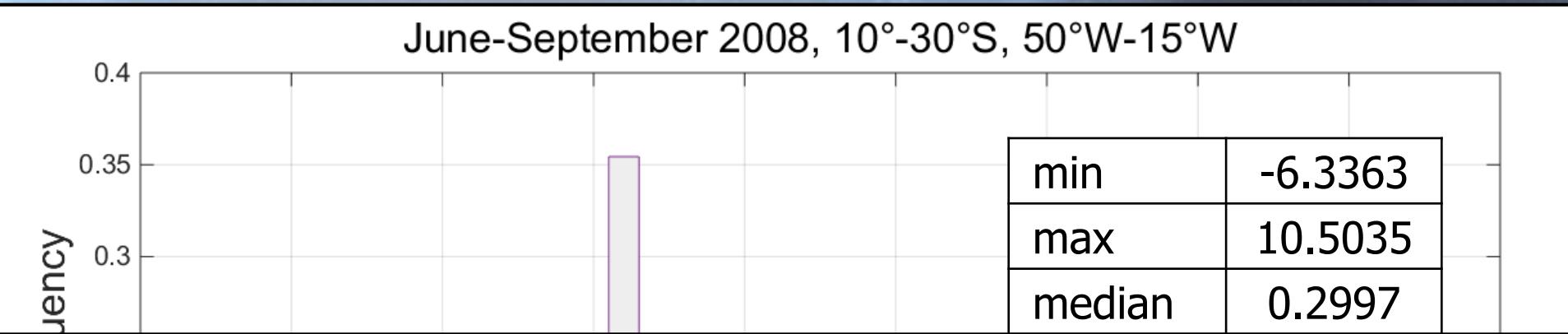
Table 6. Fine-mode lidar ratios S_λ for the optical models developed for each AERONET site. Data in parentheses are for $\tau_{f,550} = 0.5$.

Site	S_{355}	S_{532}	S_{1064}
Alta Floresta	$86.6 - 3.50\tau_{f,550}$ (84.9)	$69.1 + 2.63\tau_{f,550}$ (70.4)	$23.3 + 4.57\tau_{f,550}$ (25.6)
Bonanza Creek	$52.5 - 5.90\tau_{f,550}$ (49.6)	$59.3 - 4.65\tau_{f,550}$ (57.0)	$37.2 + 4.91\tau_{f,550}$ (39.7)
Cuiaba	$96.0 - 2.51\tau_{f,550}$ (94.8)	$67.3 + 7.09\tau_{f,550}$ (70.9)	$21.8 + 6.43\tau_{f,550}$ (25.0)
Jabiru		$73.8 + 2.39\tau_{f,550}$ (75.0)	$23.5 + 4.98\tau_{f,550}$ (26.0)
Mongu		$72.0 + 6.86\tau_{f,550}$ (75.4)	$24.8 + 4.89\tau_{f,550}$ (27.3)
Moscow		$66.6 - 0.573\tau_{f,550}$ (66.3)	$32.9 + 6.92\tau_{f,550}$ (36.3)
Mukdahan		$77.5 + 3.83\tau_{f,550}$ (79.4)	$30.6 + 8.65\tau_{f,550}$ (35.0)
Skukuza		$69.6 + 3.88\tau_{f,550}$ (71.6)	$22.2 + 4.06\tau_{f,550}$ (24.3)
Tomsk 22		$67.3 + 1.24\tau_{f,550}$ (67.9)	$31.6 + 6.25\tau_{f,550}$ (34.7)
Yakutsk	$66.6 - 6.04\tau_{f,550}$ (63.6)	$66.6 - 2.90\tau_{f,550}$ (65.1)	$29.9 + 5.80\tau_{f,550}$ (32.8)





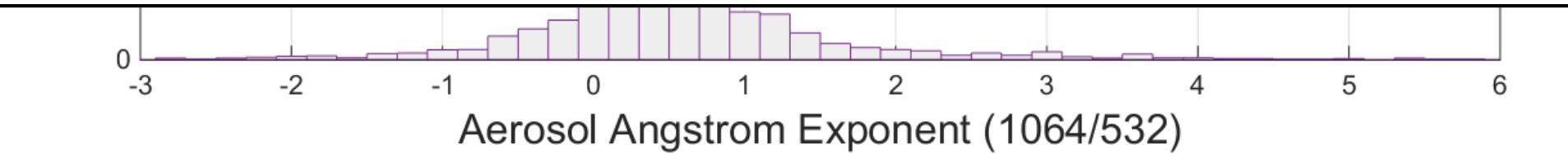
CALIOP ÅNGSTRÖM EXPONENTS FOR DUST



Freudenthaler et al., 2008: doi:10.1111/j.1600-0889.2008.00396.x

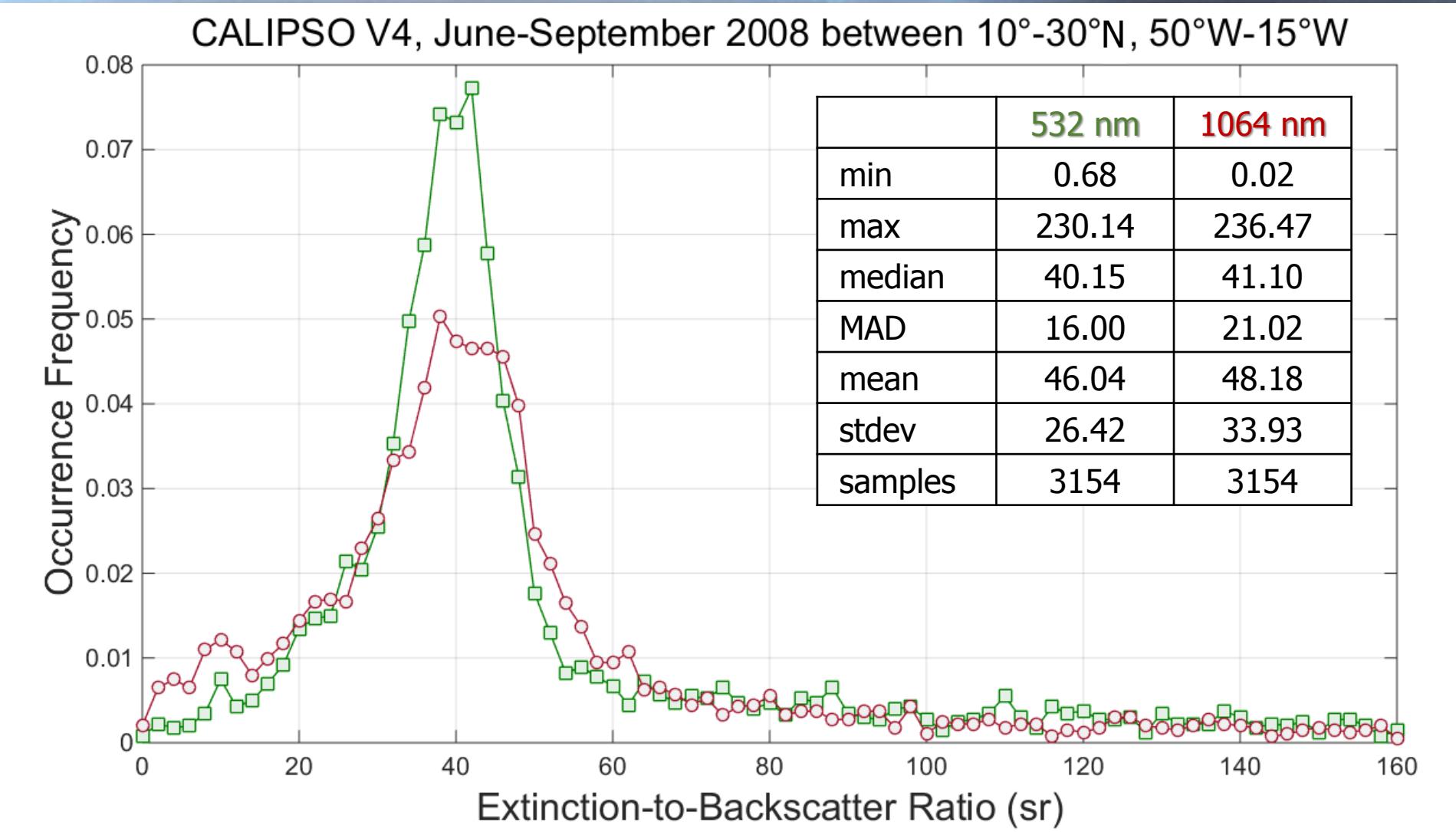
**Depolarization ratio profiling at several wavelengths in pure
Saharan dust during SAMUM 2006**

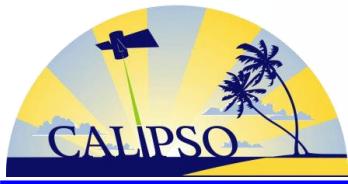
**Over the whole SAMUM period pure dust layers show [...] a mean
Ångström exponent (AE, 440–870 nm) of 0.18 (range 0.04–0.34)**





CALIOP LIDAR RATIOS FOR DUST





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thanks for listening

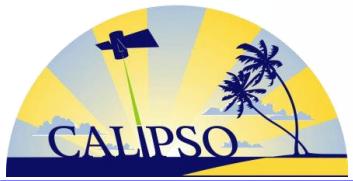
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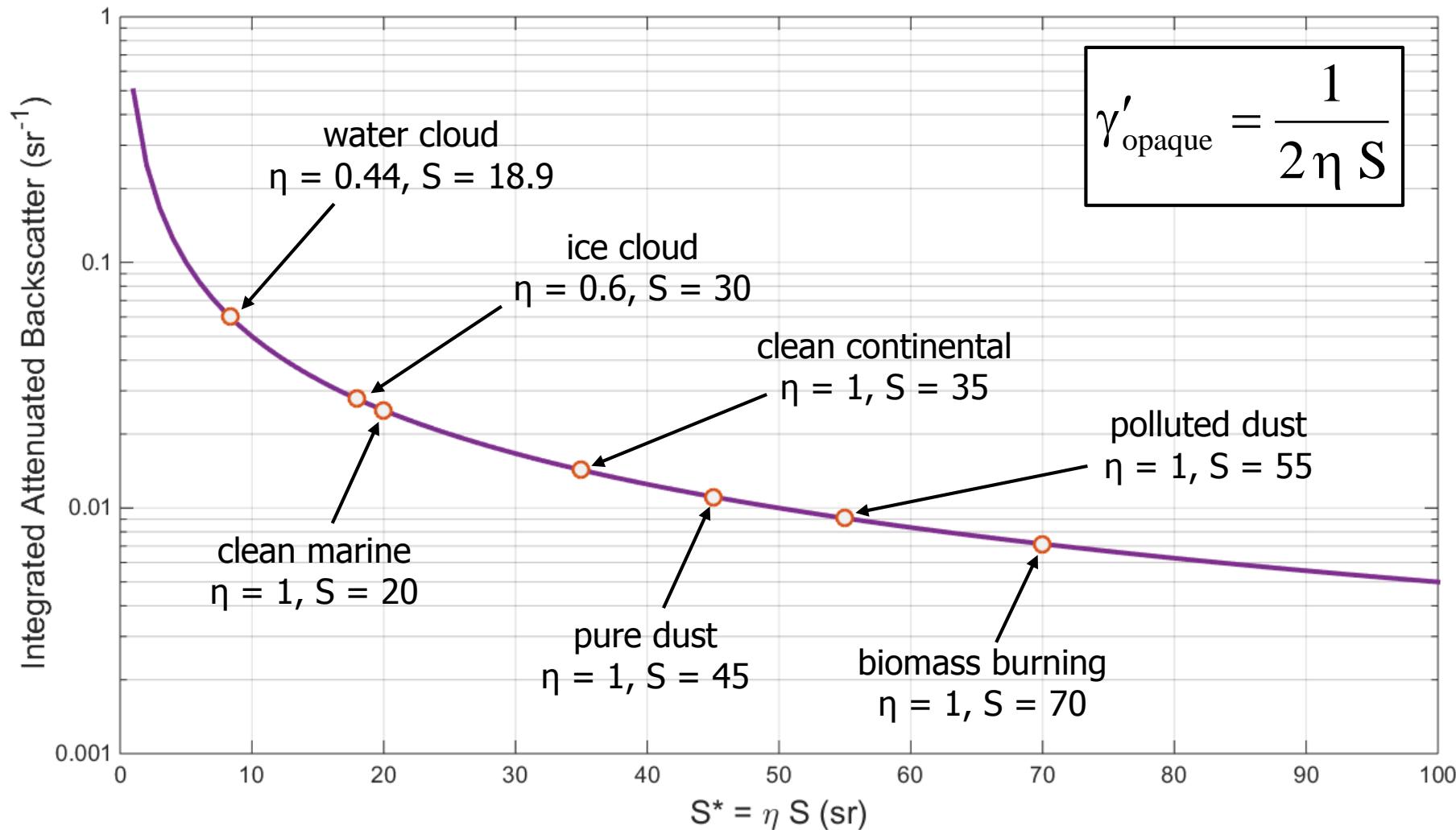


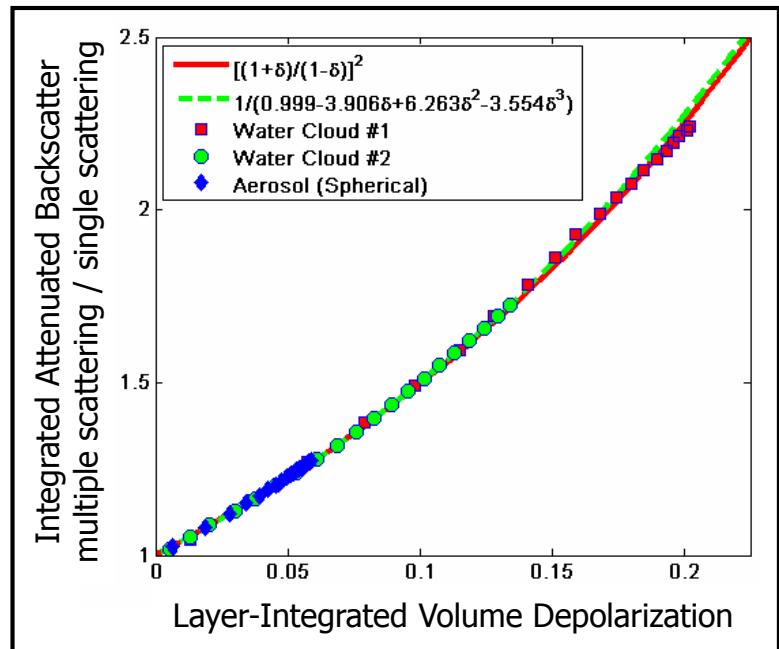
BACKUP SLIDES?





INTEGRATED ATTENUATED BACKSCATTER AS A FUNCTION OF EFFECTIVE LIDAR RATIO

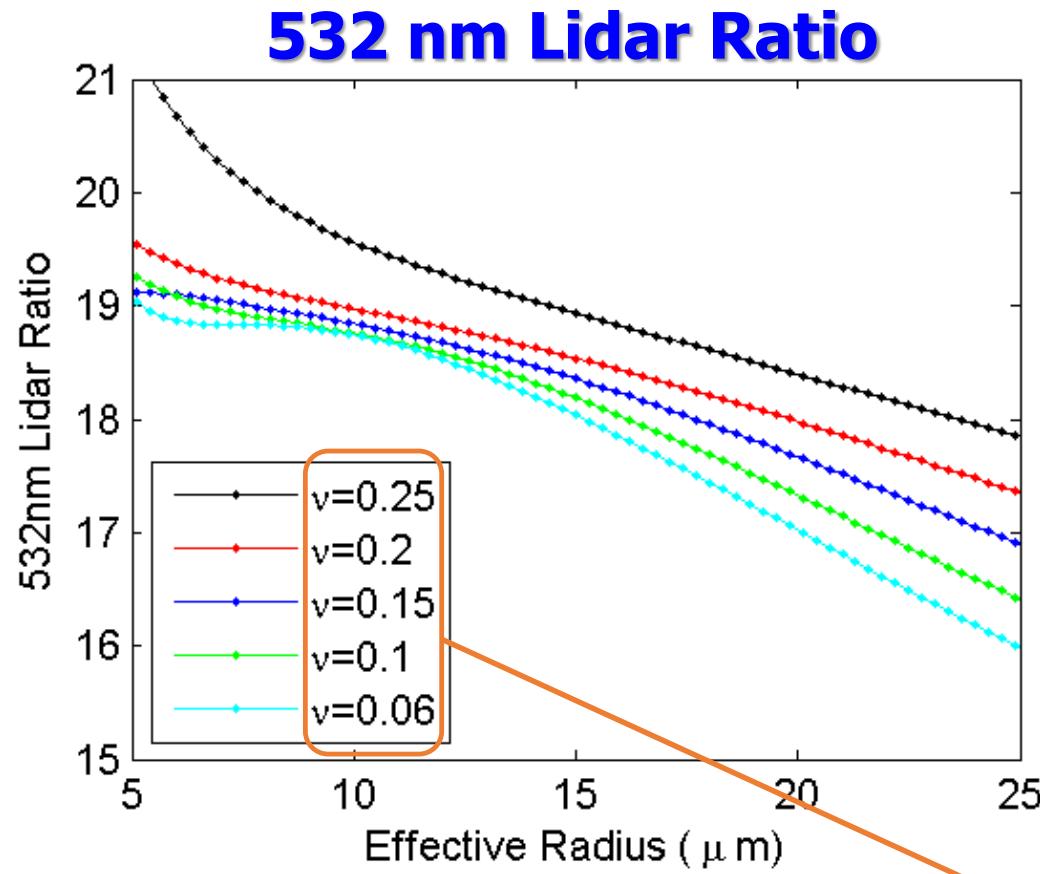




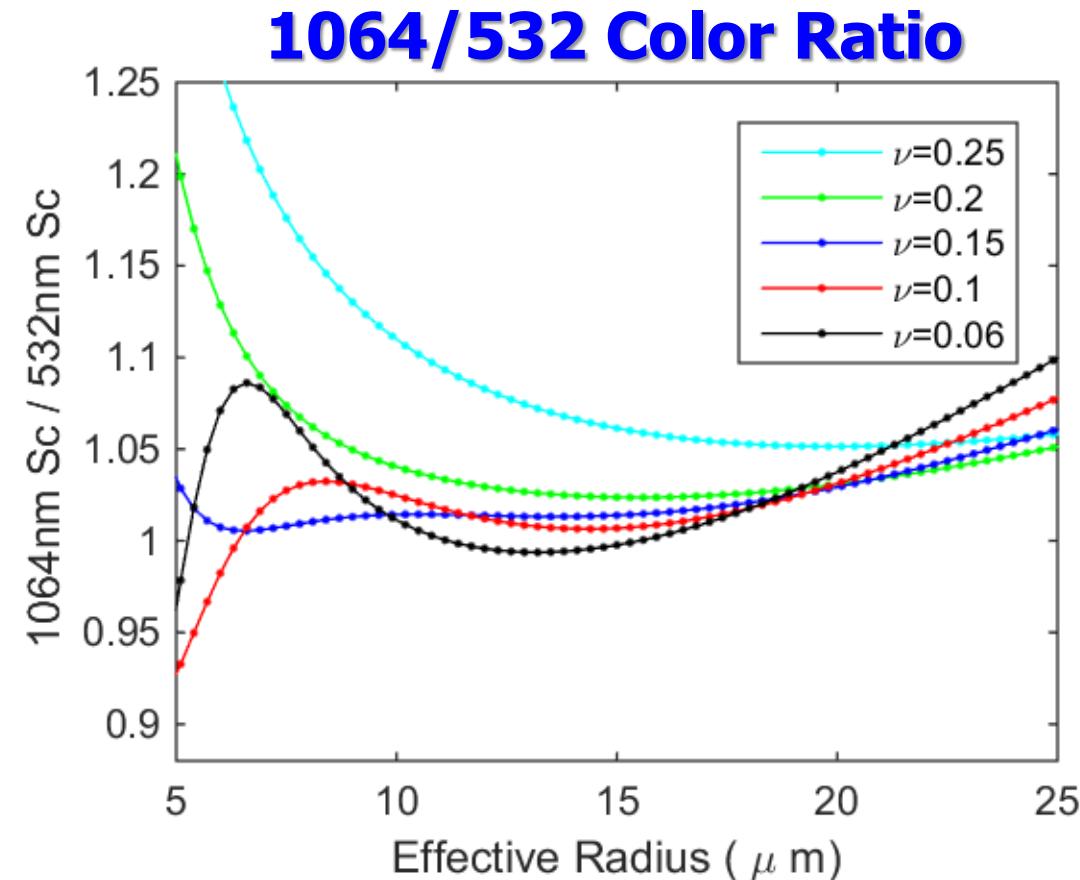
Liu et al., 2015: " S_c is found to vary insignificantly for a wide variety of water clouds, having a mean value of 18.9 sr and a standard deviation of 0.25 sr over ocean and 0.47 sr over land."

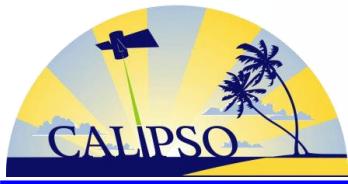


THEORETICAL CALCULATIONS OF WATER CLOUD COLOR RATIOS (SINGLE SCATTERING ONLY!)



variance of droplet size distribution





SIBERIAN SMOKE OVER NORTH AMERICA

